# LASER REMOTE SENSING II

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The problem of the analysis of lidar data is related to problems of lidar signal interpretation. Despite the wide variety of the lidar systems developed for periodical and routine atmospheric measurements, no widely accepted method of lidar data inversion or analysis has been developed or adopted. A researcher interested in the practical application of lidars soon learns the following: (1) no standard analysis method exists that can be used even for the simplest lidar measurements; (2) in the technical literature, only scattered practical recommendations can be found concerning the derivation of useful information from lidar measurements; (3) lidar data processing is, generally, considered an art rather than a routine procedure; and (4) the quality of the inverted lidar data depends dramatically on the experience and skill of the researcher.

Kovalev & Eichinger – Elastic Lidar



Solving a lidar problem is a kind of Inverse Problem









## Information Available from Aerosol Lidar

• Aerosol layer heights

- Attenuated backscatter
- Extinction profile derived from backscatter
- Extinction profile using column constraint
- Backscattering and extinction profiles at multiple wavelengths
- Particle size distribution at different heights ( so volume, surface, number densities, effective radius)
- Complex refractive index and SSA

Multiwavelength Lida

Backscatter Lidar-

Passive

Lida

Backscatter

#### FROM RAW DATA TO ETERNITY



The "non-exciting" typical lidar retrieval



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The "somehow-exciting" typical lidar retrieval









#### Summed Files for Data Processing







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#### **Summed Files for Data Processing**





#### **INSPECTING THE FILES**





#### **BACKSCATTER LIDARS**

**AEROSOL STUDIES WITH LIDARS – SOLUTIONS** 

RANGE CORRECTED SIGNAL

$$S(R) \equiv ln[R^2 P(R)]$$

RANGE CORRECTED SIGNAL





**Aerosol Scattering Ratio** 

$$\mathcal{R}(\lambda_{\rm L},\,{\rm r}) = \frac{\beta_{\pi}^{\rm mol}\left(\lambda_{\rm L},\,{\rm r}\right) + \beta_{\pi}^{\rm aer}\left(\lambda_{\rm L},\,{\rm r}\right)}{\beta_{\pi}^{\rm mol}\left(\lambda_{\rm L},\,{\rm r}\right)} = 1 + \frac{\beta_{\pi}^{\rm aer}\left(\lambda_{\rm L},\,{\rm r}\right)}{\beta_{\pi}^{\rm mol}\left(\lambda_{\rm L},\,{\rm r}\right)}$$



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#### SEPARATING MOLECULES FROM PARTICLES













## AEROSOLS AS TRACERS – ATMOPHERIC DYNAMICS



Figure 2.5: Idealized PBL daily cycle. Source: Adapted from STULL. 1988.





Figure 2.7: PBLH detected by Temperature Gradient Method from SBI and TSL height. Temperature (left), Potential Temperature (center) and Gradient of Potential Temperature (right). Source: Own author.









Figure 5.1 – Temporal evolution of RCS profile and PBLH provided by MWR (pink stars), elastic (green stars) and Doppler lidar (black stars)





**AEROSOL STUDIES WITH LIDARS – SOLUTIONS** 

#### ATMOSPHERIC OPTICAL PARAMETERS







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**AEROSOL STUDIES WITH LIDARS – SOLUTIONS** 

LIDAR RATIO

$$P(\lambda,R) = P_o \frac{A_o}{R^2} \beta_{\pi}(\lambda,R) \varepsilon(\lambda) \zeta(R) \cdot \left(\frac{c\tau_L}{2}\right) e^{-2\int_0^R \alpha(\lambda,r) dR}$$

$$L_{aer}(R) = \frac{\alpha_{mol}(R)}{\beta_{mol}(R)} = \frac{8\pi}{3}sr$$
LIDAR RATIO (MOLECULAR)

$$L_{aer}(R) = \frac{\alpha_{aer}(R)}{\beta_{aer}(R)}$$

LIDAR RATIO (AEROSOL)



#### **AEROSOL STUDIES WITH LIDARS - SOLUTIONS**



**KOVALEV & EICHINGER** 



**AEROSOL STUDIES WITH LIDARS - LIDAR EQUATION SOLUTIONS** 

$$\frac{d\chi(R)}{dR} = \frac{1}{\chi(R)} \frac{d\chi(R)}{dR} - 2\chi(R)$$

**BERNOUILLI EQUATION** 

$$\chi(R) = L_{aer}[\beta_{aer}(R) + \beta_{mol}(R)]$$

- 1. KOVALEV & EICHINGER
- 2. WEITKAMP
- 3. MEASURES



## Backscatter Lidar + Passive



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#### **AEROSOL STUDIES WITH LIDARS – AERONET CLOSURE**



#### **AEROSOL STUDIES WITH LIDARS - AERONET CLOSURE**

$$AOT_{aeronet}(\lambda) \equiv \tau(\lambda, z) = \int_{0}^{\infty} \alpha(\lambda, z) dz$$

$$AOT_{aeronet} \equiv AOT_{lidar}(\lambda) = \int_{0}^{\infty} \alpha_{aer}(\lambda, z) dz$$

$$AOT_{lidar}(\lambda) = L_{aer} \times \sum \beta_{aer}(\lambda, z) \Delta z$$



### Multiwavelength Lidar



- Joint use of  $\alpha$  and  $\beta$  key for successful retrieval.
- Minimum number of input data 5.
- The most practical configuration of Raman lidar is based on tripled Nd:YAG laser:

Three elastic channels and two nitrogen Raman (3+2) Backscattering  $\beta$  – 355, 532, 1064 nm Extinction  $\alpha$  – 355, 532 nm

**Retrieval of particle parameters** 





## Lidar Equation

$$P(z): \frac{1}{z^2}\beta(z)\exp[-2\int_0^z\alpha(z')dz']$$

Single-wavelength equation contains two unknowns: particle extinction ( $\alpha$ ) and backscattering ( $\beta$ ).

We need molecular scattering!

Raman lidar (or HSRL) can provide an additional equation.

$$P_{R}(z): \frac{1}{z^{2}}N_{R}(z)\exp\left[-\int_{0}^{z}(\alpha_{\lambda}(z')+\alpha_{\lambda_{R}}(z'))dz'\right]$$



### High Spectral Resolution Lidar (HSRL) Technique Iodine Vapor Filter Implementation)



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#### $DOI: 10.\,1175\,/\,BAMS\_D\_12\_00227.\,1$



## Retrieval of particle parameters

Requirements for retrieval algorithm

- Small number of data
- Unknown complex refractive index
- Unknown particle size distribution

We assume:

- Refractive index is size and wavelength independent;
- Particles parameters are constrained as 0.05μm<Rmin; Rmax<20 μm; 1.33<m<sub>R</sub><1.65; 0<m<sub>I</sub><0.03</li>


## **Constrained** inversion

Optical data ( $\alpha$  or  $\beta$ ) are related with PSD f(r) integral equation :

$$g_i = \int_0^\infty K_i (m, r, \lambda) f(r) dr$$

PSD f(r) is approximated by superposition of base functions  $B_i$ :





# **Constrained inversion**

Integral equations are transformed to the system of linear equations

$$g_i = \sum_{j=1}^N A_{ij} C_j.$$

Here C<sub>j</sub> – coefficients of decomposition, Aij – integrals of base functions and kernels

> Imposing constrain on solution smoothness obtain Twomey – Tikhonov expression

$$\mathbf{C} = (\mathbf{A}^{\mathrm{T}} \mathbf{A} + \boldsymbol{\gamma} \mathbf{H})^{-1} \mathbf{A}^{\mathrm{T}} \mathbf{g}$$

γ- regularization parameter determined from input data H – smoothing matrix

For every set Rmin, Rmax,  $m_R$ ,  $m_I$  - the problem is welldetermined: number of optical data = number of equations

*Rmin, Rmax, m<sub>R</sub>, m<sub>I</sub> are varied and several thousands of solutions are found* 

Every solution is characterized by discrepancy (difference between experimental data and data calculated from solution)

Solutions are ranged and averaged



Particles parameters are retrieved from synthetic optical data; ~1% of solutions are averaged.



# **Numerical simulation**

Model PSD



-

Introducing errors



- Synthetic input optical data were calculated from bimodal PSDs;
- Errors in  $[0, \varepsilon]$  range were introduced in input data in a random way;

Uncertainty of particle parameters retrieval for 10% input errors:

- Effective radius, volume, surface density -~30% accuracy
- Real part of refractive index -±0.05
- Imaginary part of refractive index 50% for mI>0.005





### Comparison of AERONET and lidar data



Parameter		AERONET	Lidar	
AOT	355	0.38	0.41±0.08	
	532	0.27	0.31±0.06	
V <sup>tot</sup> (μn	ո <sup>3</sup> /µm²)	0.073	(0.076±0.015)	
V <sup>fine</sup> (μm³/μm²)		0.05	(0.057±0.012)	
r <sup>tot</sup> <sub>eff</sub> (μm)		0.3	0.33±0.07	
r <sup>fine</sup> <sub>eff</sub> (μm)		0.214 0.25±0.05		
m <sub>R</sub>		1.37	1.34±0.05	
m		0.006	0.009±0.0045	

•Time lag between lidar and AERONET

measurements is about 2 hours

- We have to extrapolate profiles below
- 1 km



# **DISCOVER-AQ 2011 CAMPAIGN**

The P-3B aircraft carried a suite of nine scientific instruments including NASA LARC HSRL system.

14 flights occurred during June 27 - July 31 period over Baltimore – Washington area.

GSFC multiwavelength Raman lidar performed the measurements from the ground

*Lidar parameters:* Telescope aperture 400 mm Laser power at 355 nm – 20 W

Operational wavelengths: elastic – 355, 532, 1064 nm Raman – 387, 608, 408 nm

- Temporal resolution of the measurements 2 min.
- Measurements are vertical, height resolution 15 m.
- Sonde data are from Beltsville (~5 miles apart)





# 20 – 21 July measurements



Aerosol is represented mainly by sulfates and biomass burning products. RH is insufficient for significant hygroscopic growth.



# Retrieval of height – temporal distributions of particle parameters













### **Time-series of particle parameters**

Particle density, effective radius and real part of CRI are shown for 1100 and 2500 m layers



Volume and number density present strong variations while effective radius is quite stable

# Lidar derived column volume



Volume profiles are extrapolated as constant below 1000 m. For comparison AERONET results at 23:00 UTC are shown.



### 21 – 22 July measurements





# Height temporal distributions of particle parameters











### **Comparison with AERONET**



### How many wavelengths do we need?

- $1\lambda$  + AERONET Particle concentration
- $2\lambda$  + AERONET Profiles of effective radius and concentration (LIRIC)
- **3**λ For known refractive index particle radii and concentration
- 3+1 For many cases radii, concentration, RI
- 3+2 Particle size distribution, RI
- 3+2+1 $\delta$  Treatment of dust mixtures
- $3+2+3\delta$  Aerosol classification



### **NETWORKS**









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- Consolidate the measurement and data acquisition protocols
- Establish a QA/QC routine among all stations
- Improve and establish an unified data analysis routine common to all stations, e.g., Single Calculus Chain
- Create a scientifically significant distributed database,e.g., lidar ratio, particle extinction, backscatter, angstrom exponents and particle depol. regional values that can be assimilated to air quality & forecast models and validation missions.









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ST.	ID	LAT LON ALT(m)	Detected Channels (nm)
Buenos Aires	VMA	-34.56° -58.51° 10	1064, 532 <sup>p</sup> & 355 <sup>p</sup>
Buenos Aires	SMN	-34.56° -58.42° 10	1064, 607, 532 <sup>p</sup> , 387 & 355 <sup>p</sup>
Neuquen	NQN	-38.95° -68.14°, 266	1064, 532 <sup>p</sup> , 266& 355 <sup>p</sup>
Bari- loche	BRC	-41.15° 71.16° 837	1064, 607, 532, 387 & 355
Com- modoro	CDR	-45.79° -67.46° 48	1064, 532 & 355
Rio Gallegos	RGL	$-51.60^{o}$ $-69.32^{o}$ 20	355, 308 & 355
Rio Gallegos II	<u>SRG</u>	-51.61° -69.31° 17	1064, 607, 532 <sup>p</sup> , 387 & 355 <sup>p</sup>
Cordoba - HRSL	COR	-31.68° -63.87° 322	1064, 607, 532 <sup>p</sup> , 408, 387 & 355 <sup>p</sup>
Punta Arenas	PAR	-53.22° -70.88° 15	1064, 607, 532 <sup>p</sup> , 408, 387 & 355 <sup>p</sup>
Tucuman	TUC	$-26.79^{o}$ $-65.21^{o}$ 485	1064, 607, 532 <sup>p</sup> , 408, 387 & 355 <sup>p</sup>

ST.	ID	LAT LON ALT(m)	Detected Channels (nm)		
S. Paulo	SPU	-23.56° -46.74° 740	1064, 607, 532, 531, 408, 387 & 355		
S. Paulo	SPT	Trans- portable	607, 532		
Manaus	MAO	-02.89° -59.97° 30	408, 387, 355		
Natal	NAT	$-05.82^{o}$ $-35.20^{o}$ 12	1064, 532 <sup>p</sup> & 355 <sup>p</sup>		
Тетисо	TMU	-38.73° -72.60° 108	532		
Medellin	MED	+06.22° -75.57° 1545	1064, 532 & 355		
Medellin CIBioFi	MEC	+03.37° -76.53° 982	1064, 532 & 355		
Medellin SIATA	MES	+03.37° -76.53° 1538	355 <sup>p</sup>		
La Paz	LPZ	16.53° 72.07° 3500	532		

- Tropospheric studies
  - PBL height determination
  - Aerosol
  - Cloud studies
- Stratospheric studies
  - Volcanic eruptions
  - Ozone
- Satellite Validation
- Outreach
- TSM studies







# lalinet-loci complex terrain





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# lalinet-troposphere: pbl studies







# lalinet-troposphere: aerosol typing



#### SP:

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# lalinet-troposphere: cirrus cloud forcing







# lalinet-stratosphere: capability extension

Scientific goals:
Establishing standardized, aerosol backscattering measurements in UTLS.
Production of standardized processing algorithms, measurement methodologies and its quality control.
Look for funds for new lidars (UTLS) approach:
Evaluate the existing lidar capabilities .

• Determine alternatives without technological changes.

Antuña-Marrero, J. C., E. Landulfo, et al., 2018: One step further in the objectives of LALINET: preparation for the next major volcanic eruption & validations of the UTLS aerosols measurements from EarthCare and Sage III satellite missions. Poster, *Chapman Conference on Stratospheric aerosol in the post-Pinatubo era: Processes, Interactions and Importance,* 18-23 March, Puerto de la Cruz, Tenerife, Spain.



# alinet-stratosphere: volcanic eruptions



Location	Layer Height (km)	AOD at 532 nm	LR (532nm)	LR (355nm)	Ångström 355/532	Туре
Leipzig	2.6-4.3	0.35	$60 \pm 5$	$60 \pm 5$	$0.03\pm0.40$	Ash
Munich	2.6-3.5	-	50-60	$50 \pm 5$	$-0.11\pm0.18$	Pure dry ash
Potenza	2.0-3.0	-	$42 \pm 2$	$50 \pm 3$	$1.4 \pm 0.2$	Sulfates with some ash
Evora	2.7-3.7	0.07	$39 \pm 2$	$32 \pm 4$	$0.68 \pm 0.63$	Fresh volcanic particles
Granada	2.6-2.9	-	$47 \pm 7$	$48 \pm 16$	$0.066 \pm 0.005$	-
Cabauw	2.7-6.0	0.53	$42 \pm 1$	$44 \pm 24$	$0.30 \pm 0.03$	Sulfate-ash mixture
Athens	3.0-4.8	0.05	$67 \pm 13$	$89 \pm 3$	$0.57 \pm 0.26$	Aged ash/sulfates
Athens	2.5-3.0	0.04	$76 \pm 5$	$78 \pm 3$	$1.72 \pm 0.06$	Sulfates
Brazil	18-19.3	0.16	$76 \pm 27$	$63 \pm 21$	$0.61 \pm 0.58$	Sulfates with some ash

#### Synergetic Aerosol Layer Observation After the 2015 Calbuco Volcanic Eruption Event

Fábio J. S. Lopes <sup>1,\*</sup> <sup>(5)</sup>, Jonatan João Silva <sup>1,2</sup>, Juan Carlos Antuña Marrero <sup>3</sup>, Ghassan Taha <sup>4</sup> and Eduardo Landulfo <sup>1</sup> <sup>(5)</sup>

Remote Sens. 2019, 11, 195; doi:10.3390/rs11020195



#### The OMPS Limb Profiler Analysis



#### Landulfo et al.



## lalinet: satellite validation

#### LALINET AEOLUS CAL/VAL







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# Lalinet satellite validation

- SPU Lidar station AEOLUS Validation –
- 66 overpasses since 04th November 2018
- 2 overpasses by week Sundays 08:49 and 21:29 UTC
- Horizontal distance from SPU Station ~ 102 km
- SPU Lidar Station:
  - 20 correlative measurements since 04th November 2018 30% of the overpasses
- Access L2A data VirES Aeolus Service https://aeolus.services/
- L2A data products: SCA Standard Correct Algorithm Backsc. and extinction profiles







### lalinet: satellite validation



Landulfo et al.

0.05 0.10 Lat: -22.74 Long: -47.36

0.00 0.01

SCA\_Backscatter (10<sup>-6</sup>m<sup>-1</sup>)

0.02

-0.01





### lalinet: outreach

WLMLA (Year)	Place	Attendees				Papers	
		LA(%)	RW(%)	Total	ST (%)	РО	OR
I (2001)	Camagüey, <b>Cuba</b>	9 (39)	14 (61)	23	5 (22)	5	14
II (2003)	Camagüey, <b>Cuba</b>	13 (52)	12 (48)	25	13 (52)	2	25
III (2005)	Popayán, <b>Colombia</b>	41 (79)	11 (21)	52	26 (50)	6	25
IV (2007)	Ilhabela, <b>Brazil</b>	30 (71)	12 (29)	42	20 (48)	16	29
V (2009)	Buenos Aires, <b>Argentina</b>	42 (65)	23 (35)	65	21 (32)	31	31
VI (2011)	La Paz, <b>Bolivia</b>	52 (81)	12 (19)	64	32 (50)	15	21
VII (2013)	Pucón, <b>Chile</b>	35 (76)	11 (24)	46	19 (41)	20	24
VIII (2015)	Cayo-Coco, <b>Cuba</b>	29 (71)	12 (29)	41	15 (37)	25	19
IX (2016)	Santos, <b>Brazil</b>	41 (84)	8 (16)	49	20 (41)	29	23
X(2018)	Medellin, Colombia	49(89)	6(11)	55	35(63)	23	13









# lalinet: outreach

Assessment of atmospheric optical Properties during biomass burning

Events and

Lidar Network

Long-range transport of desert dust

Building a bridge between EARLINET and LALINET

ALCANTARA EUROPEAN SPACE AGENCY -

INOE(Rumania), LMD (Germany) &

UGR (Spain)

SPU + MAO + NAT

- 1. Student Training Lical
- 2. QC/QA @ LALINET sites
- 3. Coordinated Campaigns
- 4. SCC x LALINET Algorithms















SCC data data comparison with MILGRAU







### lalinet: outreach







#### SPSAS on Atmospheric Aerosols


## lalinet: acknowledgements

